Tievebaun GWB: Summary of Initial Cha	racterisation.
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Hydrometric Area Local Authority Hydrometric Area 36 Leitrim Co. Co. Sligo Co. Co.		Associated surface water bodies	Associated terrestrial ecosystems	Area (km²)		
		<i>Rivers:</i> None <i>Streams:</i> 17 unnamed streams <i>Lakes:</i> None	Ben Bulben, Gleniff and Glenade Complex (O'Riain, 2004)	10		
	Topography	of Truskmore (Figure 1). The southern bound the remaining boundaries consisting of low	l land GWB covering Tievebaun Mountain and t lary comprise a topographic divide (Hydromet er permeability rocks. Being an essentially u from 360 to 610 mAOD on Tievebaun, an pe although very few channels are mapped.	ric Area 35), with upland GWB, th		
	Aquifer type(s)	Rk ^e : Regionally important karst aquifer dominated by conduit flow, covers the majority of the GWB (80%). The areas are capped by Lm : Locally important aquifer which is generally moderately productive (11%), which is surrounded by a thin band of LI : Locally important aquifer which is moderately productive only in local zones.				
	Main aquifer lithologies	The GWB mainly comprises three rock groups of Dinantian age: Pure Bedded Limestones (79.57%), Sandstones over the upland caps (11.29%) and an intervening band of Mixed Sandstones Shales and Limestones (9.14%). Refer to Table 1 for full details.				
	Key structures	The rocks are generally dipping towards Tieve	baun peak by 10-40°.			
Aquifers	Key properties	No abstraction/discharge data are available for this GWB however, highly karstified aquifers are often associated with extremely variable transmissivity values, borehole yields and spring yields. Spring yields can also be very large. Recharge can be rapid and a large proportion of the flow can occur through conduits, sometimes at extremely high velocities (e.g 100s m/hr). Accordingly highly karstified rocks are often associated with low storativity.				
Geology and Aquifers		No karst data are available although the c.20 features were recorded in the nearly Glenaniff GWB, wh comprises the same rocks in a similar sequence. The location of these features indicate that the aquifer is hig karstified. This is likely to be case for the Tievebaun GWB, especially as there is very low drainage density o the rocks, suggesting that water rapidly recharges the aquifer, which is characteristic of karst rocks. Groundwater gradients cannot be calculated although flow directions are likely to radiate downslope to north, west and east within the GWB. Regional flow is generally northwards, to eventually discharge in Donegal Bay.				
	Thickness	(Pure Bedded Limestones Aquifer Chapter) In the pure limestones, most groundwater flows in an epikarstic layer a couple of metres thick and in a zone of interconnected solutionally-enlarged fissures and conduits that extends approximately 30 m below this. Most groundwater flux in the Sandstones (Lm aquifer) is also likely to be in uppermost top 30 m (c.3 m broken, weathered material underlain by interconnected fissuring), although there will also be a zone of isolated, poorly connected fissures – typically less than 150 m bgl – in all of the rock types. In the lower permeability Mixed Sandstones, Shales and Limestones, there is a lower potential for deeper flows and the more interconnected fissure zones is also likely to be shallower –c.10-30 m in thickness.				
ta	Lithologies	No data are available for 60% of this GWB (NI and Leitrim). However, of the remaining area, till is the predominant subsoil (c.70%), with a small proportion of peat (12%).				
Overlying Stra	Thickness % area aquifer	Although available data are limited, the distribution of outcrops and steeper topography suggest that the subsoil in this GWB is relatively thin i.e. mainly less that 3 m thick. <i>[Information will be added at a later date]</i>				
arlyi.	near surface	[injormation will be added at a tater date]				
õ	Vulnerability	Although vulnerability data are not available, vulnerability is <i>likely</i> to be Extreme over the majority of the GWB, based on the assumption that the subsoil is generally thinner.				
Recharge	Main recharge mechanisms	Both point and diffuse recharge occur in this GWB. Diffuse recharge occurs via rainfall percolating through thin subsoil and outcrops over all of the aquifer types. In the pure limestones, point recharge to the underlying aquifer occurs via any swallow holes, caves and dolines. Although recharge along 'losing' sections of streams is also associated with this particular type of karst aquifer, to date none have been recorded in this GWB. The steeper slopes in the body will promote runoff, however, the permeability nature of both the pure limestones and the sandstones are expected to results in a high proportion of recharge occurring over the majority of the body. This is also indicated by the low stream density.				
	Est. recharge rates	[Information will be added at a later date]				

	Important	Sources: None identified.			
	springs and high vielding wells	Springs: None identified.			
	yielding wens	Excellent Wells: None identified.			
		Good Wells: None identified.			
Discharge	Main discharge mechanisms	Any groundwater discharges are to the streams, rivers and any springs found within the body. However, given the permeable nature of the aquifers and the topographic location of this GWB, there are likely to be minimal discharges within the GWB itself. Instead, a large proportion of groundwater moving down-gradient through the permeable limestone is more likely to discharge to surface water once it encounters the lower permeability aquifers of the Rossinver GWB. This assessment of discharge is indicated by the surface drainage pattern; the majority of rivers/streams in the area start at the boundary between the Tievbaun and Rossinver GWBs.			
scha	Hydrochemical	National classification: Dinantian Pure and Impure Limestones			
Di	Signature	Calcareous. Generally Ca-HCO ₃ signature. Alkalinity (mg/l as CaCO ₃): range of 10-990; mean of 283 (2454 data points) Total Hardness (mg/l): range of 10-1940; mean of 339 (2146 data points) Conductivity (μ S/cm): range of 76-2999; mean of 691 (2663 data points)			
		<i>National classification:</i> Dinantian Sandstones Calcareous. Generally Ca-HCO ₃ signature. Alkalinity (mg/l as CaCO ₃): range of 5-524; mean of 153 (65 'non limestone subsoils' data points) Total Hardness (mg/l): range of 5-502; mean of 162 (67 'non limestone subsoils' data points) Conductivity (μ S/cm): range of 39-1184; mean of 408 (69 'non limestone subsoils' data points)			
		(Calcareous/Non calcareous classification of bedrock in the Republic of Ireland report)			
Groundwater Flow Paths		As these rocks are generally devoid of inter-granular permeability, groundwater flows through fissures, faults, joints and bedding planes. In pure bedded limestones, these openings are frequently enlarged by karstification resulting in significantly enhanced rock permeability. Karstification can be also accentuated along structural features such as fold axes and faults. An epikarst layer in the upper few metres of the rock is likely to be present on top of the karstified aquifer. Shallow groundwater flow is likely to be dominant, although a component of deep groundwater flow would be expected.			
		Groundwater flow through karst areas is frequently extremely complex and difficult to predict. Although minimal hydrogeological investigation has been undertaken in this particular region, karst aquifers that are dominated by conduit flow are generally capable of rapidly transmitting large volumes of groundwater, which is frequently localised. Flow velocities are known to be variable, both temporally and spatially, with groundwater flows often exhibiting a rapid response to rainfall events, giving rise to 'spikey' hydrographs and springs with highly variable discharge. Rapid, localised flow through conduits often results in the aquifer having low storativity. Groundwater flow directions deviating from the expected (i.e. perpendicular to the assumed water table contours), and in extreme cases c) flow across surface catchment divides/ beneath surface channels. This, however, depends on the frequency of faults, fissures and joints, and has not been established for this GWB.			
		Groundwater flow is thought to be mainly unconfined. In the karstified aquifers, groundwater flow is regional scale – flow path lengths of several kilometres are not unusual, which would span the width of the GWB. Long flow paths are also likely be occur in the Sandstones although shorter flow paths are associated with the Mixed Sandstones, Shales and Limestones/ Impure Limestones (c.100-300 m). Overall, groundwater flow is likely radiate out from the summits in the GWB. However, the karstified nature of the pure limestone means that locally, groundwater flow directions can be highly variable.			
surface water sufface water sufface water sufface water interactions 'g		There is a high degree of interconnection between groundwater and surface water in karstified limestone areas such as in this GWB. Any unrecorded swallow holes, dolines, caves, turloughs, springs, and 'losing' and 'gaining' streams all provide a direct route between surface water and groundwater systems. This rapid interchange between surface water and groundwater is often reflected in their similar water quality as contamination is also rapidly transported between the two systems.			

		Northern, western and eastern boundaries comprise lower permeability aquifers. A topographic divide provide the southern boundary. Essentially constituting an upland area, the topography is generally mountainous.					
	Appro	The main rock type in this GWB (c.80%) is a karstified limestone that is dominated by conduit groundwater flow (\mathbf{Rk}^{c} Approximately 10% of the area is underlain by sandstones, which are also considered to be characterised by a productive fracture flow system (Lm). The Sandstones are ringed by Mixed Sandstones, Shales and Limestone (Ll).					
	broker limest	tost of the unconfined groundwater flux is in the uppermost 30 m of the aquifers. This occurs through a few metres (c.3 m) or oken, weathered bedrock and an underlying zones of interconnected joints, fissures, fractures and faults. In the pure nestones, the upper weathered zone is likely to equate to an epikarst layer and the underling joints, fissures, fractures and ults will be karstified (solutionally enlarged). Deeper groundwater flow may occur along permeable fault or fracture zones.					
ll model	pure li	Transmissivity values and well/spring yields are likely to be variable, reflecting zones of higher and lower permeability. In the pure limestones, there is the potential for high flow velocities of large quantities of groundwater through conduits, coupled with low storativity, resulting in rapid flow/discharge responses to rainfall.					
Conceptual model	paths (general, the degree of interconnection in karstic systems is high and they support regional scale flow systems. Long flow the (kilometres in length) can be expected although are likely to be shorter in discharge areas (100-300 m). Similar flow path to get swould be expected in the sandstones.					
ŭ	Recha	rge occurs by:					
		liffuse means in all rock types - via outcrops and through thin subsoil, and					
		additional point mechanisms in the karstified limestones; swallow holes, dolines, caves and along lengths of losing streams – mainly occurring where subsoils are thin i.e. areas of extreme vulnerability.					
		Due to the combination of point recharge and rapid flow through solutionally enlarged joint/fissure/fracture zones, there is minimal potential for contaminant attenuation in the limestone aquifer.					
	throug	• The main discharges are to the rivers and springs, principally at the boundary of the GWB, where the groundwater moving through the karstified aquifer encounters the lower permeability aquifers of the Rossinver GWB. The flow radiates out from the centre of the GWB to the west, north and east, as determined by the topography.					
		is a high degree of interaction between surface water and groundwater in this GWB.					
Attacl	hments	Figure 1. Table 1.					
Instru	umentation	Stream gauges: None identified. EPA Water Level Monitoring boreholes: None identified. EPA Representative Monitoring points: None identified.					
Information Sources		 El A Representative Monitoring points, None identified. MacDermot, C.V. Long C.B. and Harney S.J (1996) Geology of Sligo-Leitrim: A geological description of Sligo, Leitrim and adjoining parts of Cavan, Fermanagh, Mayo and Roscommon, to accompany bedrock geology 1:100,000 scale map, Sheet 7, Sligo - Leitrim. With contributions from K. Carlingbold, G. Stanley, D. Daly and R. Meehan. Geological Survey of Ireland, 100pp. 					
		O' Riain, 2004. <i>Water Dependent Ecosystems and Subtypes (Draft)</i> . Compass Informatics in association with National Parks and Wildlife (DEHLG). WFD support projects.					
	aimer	Note that all calculation and interpretations presented in this report represent estimations based on the information					

Figure 1. Location and Boundaries of GWB

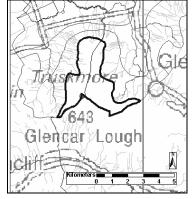


Table 1. List of Rock units in Tievebaun GWB

	Rock Unit Name	Code	Description	Rock Unit Group	Aquifer Class.	% Area
]	Dartry Limestone Formation	DA	Dark fine-grained cherty limestone	Dinantian Pure Bedded Limestones	Rkc	79.57%
(Glenade Sandstone Formation	GD	Pale orthoquartzitic sandstone	Dinantian Sandstones	Lm	11.29%
				Dinantian Mixed Sandstones, Shales and		
]	Meenymore Formation	ME	Shale, laminated carbonate, evaporite	Limestones	Ll	9.14%